Geochemical evidences of trace metal anomalies for finding hydrocarbon microseepage in the petroliferous regions of the Tatipaka and Pasarlapudi areas of Krishna Godavari Basin, India

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Abstract: The long-term seepage of hydrocarbons, either as macroseepage or microseepage, can set up near-surface oxidation reduction zones that favor the development of a diverse array of chemical and mineralogical changes. The bacterial oxidation of light hydrocarbons can directly or indirectly bring about significant changes in the values of pH and Eh of the surrounding environment, thereby also changing the stability fields of the different mineral species present in that environment. The paper reports the role of hydrocarbon microseepage in surface alterations of trace metal concentrations. In this study trace metal alterations were mapped that appear to be associated with hydrocarbon microseepages in the oil/gas fields. A total of 50 soil samples were collected near oil and gas fields of the Tatipaka and Pasarlapudi areas of the Krishna Godavari Basin, Andhra Pradesh. The soil samples were collected from a depth of 2-2.5 m. The paper reports the chemical alterations associated with trace metals in soils that are related to hydrocarbon microseepages above some of the major oil and gas fields of this petroliferous region. Trace metals, such as scandium (Sc), vanadium (V), chromium (Cr), cobalt (Co), nickel (Ni), copper (Cu), zinc (Zn), barium (Ba) and strontium (Sr), in soil samples were analyzed using inductively coupled plasma-mass spectrometry (ICP-MS). The concentrations of Sc (8 to 40 mg/kg), V (197 to 489 mg/kg), Cr (106 to 287 mg/kg), Co (31 to 52 mg/kg), Ni (65 to 110 mg/kg), Cu (88 to 131 mg/kg), Zn (88 to 471 mg/kg), Ba (263 to 3,091 mg/kg) and Sr (119 to 218 mg/kg) were obtained. It was observed that the concentrations of trace elements were tremendously increased when they were compared with their normal concentrations in soils. The analysis of adsorbed soil gas showed the presence of high concentrations of ΣC_{2+} (C_2H_6, C_3H_8 and n-C_4H_{10}) ranging from 7 to 222 μg/kg respectively. Integrated studies of trace elements over adsorbed light gaseous hydrocarbons (ΣC_{2+}) anomalies showed good correlation with the existing oil and gas wells. The carbon isotopic composition of $\delta^{13}C_1$ of the samples ranges between –36.6‰ to –22.7‰ (Pee Dee Belemnite) values indicate thermogenic origin, which presents convincing evidence that the adsorbed soil gases collected from these sediments are of catagenetic origin. The increase in the concentrations of trace metals near oil/gas producing areas, suggests a soil chemical change to a reducing environment, presumably due to the influence of hydrocarbon microseepage, which could be applied with other geo-scientific data to identify areas of future hydrocarbon exploration in frontier areas.

Key words: Hydrocarbons, microseepage, trace metal alterations, adsorbed soil gas, Krishna Godavari Basin

1 Introduction

Surface geochemical prospecting for hydrocarbons comprises investigation of near surface soils/sediments for occurrence of hydrocarbons that may indicate the location of subsurface petroleum reservoirs. Surface geochemical

methods are based on the premise that the component hydrocarbons gases (CH_4, C_2H_6, C_3H_8 and n-C_4H_{10}) migrate to the surface from the sub-surface petroleum accumulations through faults and fractures and leave their signatures in the near surface soils. Many hydrocarbon migration mechanisms such as diffusion, effusion, advection with moving waters and permeation have been proposed and studied by various workers (Price, 1986; Tedesco, 1995; Schumacher and Abrams, 1996). These light gaseous hydrocarbons migrate

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vertically to the surface and get adsorbed in the near surface soil matrix (Jones and Drozd, 1983; Klusman, 1993; Tedesco, 1995; Kalpana et al, 2010; Schumacher and LeSchack, 2002). These light gaseous hydrocarbons may be detected either directly (adsorbed soil gas or free gas) or indirectly (microbial indicators, trace element geochemistry) through the geochemical changes they induce and an anomaly at the surface can be reliably related to a petroleum accumulation at a deeper level. The long-term seepage of hydrocarbons, either as macroseepage or microseepage, can set up near-surface oxidation reduction zones that favor the development of a diverse array of chemical and mineralogical changes (Donovan, 1974; Petrovic et al, 2008; Khan and Jacobson, 2008). Some bacteria utilize hydrocarbon gases as their sole source of food and are found to be enriched in the near surface soils/sediments above oil and gas reservoirs, hence the bacteria are used as indicators for finding oil and gas reservoirs (Rasheed et al, 2011). The bacterial oxidation of light hydrocarbons can directly or indirectly bring about significant changes in the pH and Eh of the surrounding environment, thereby also changing the stability fields of the different mineral species present in that environment. These changes result in the precipitation or dissolution and remobilization of various mineral species and elements, such that the rock column above a leaking petroleum accumulation becomes significantly and measurably different from laterally equivalent rocks (Pirson, 1969; Oehler and Sternberg, 1984; Price, 1986). It is also believed that hydrocarbons along with inorganic components associated with oil are continuously migrating to the surface which forms the basis for surficial petroleum investigations with regard to trace metals (Petrovic et al, 2008). The alterations occur because of leaking hydrocarbon. Staunton (2002) has reviewed some of the interactions of metals with soil organic matter. The effect of organic matter may be by interaction due to their adsorption properties and complexation capacity, or indirectly because of the modification of mineral surface properties with organo-mineral complexes being formed. Trace elements in soils have various origins which include lithogenic elements that are directly inherited from the lithosphere (parent material); anthropogenic elements are direct or indirect results of human activities and pedogenic elements are of lithogenic and anthropogenic origins but their distribution in soil horizons and soil particles are changed due to mineral transformation and other pedogenic processes. Soils consists of a heterogeneous mixture of different organic and organo-mineral substances, clay minerals, (hydrous) oxides of Al, Fe and Mn and other solid components, as well as a variety of soluble substances. The binding mechanisms for trace metals are therefore complex and vary with the composition of the soil, soil acidity and redox conditions. The complexity of all possible reactions occurring in soil includes dissolution, sorption, complexation, migration, precipitation, occlusion, diffusion, binding by organic substances, absorption & sorption by microbiota and volatilization. All these processes are governed by several soil properties of which pH and redox potential are known to be the most important and any changes in these parameters affect most of the processes regulating

the speciation of any metal in soil (Kabata-Pendias, 1993; Calmano et al, 1993). The metal ions will be adsorbed by soils and their constituents by several mechanisms like non-specific cation-exchange, specific adsorption, co-precipitation and organic complexation (Gast, 1979). The metal ions will also be adsorbed on the surface of the soil colloids like humus polymers, clay minerals and oxides.

Recent research in soil science has revealed that changes in pH and Eh have an enormous effect on chemical and biological processes in the soil substrate. Surface geochemical surveys typically do not measure Eh and pH as standard procedure. Petroleum microseepage into the soil substrate causes several chemical reactions and microbial oxidation of hydrocarbons and also causes decreases in Eh and pH. Mineral stability in any environment is dependent on, and is a function of pH and Eh. Hydrocarbons, migrating into the near-surface soils destabilizes many compounds and increases the solubility of the trace and minor elements. Many metal elements are highly soluble in acid solutions but will be precipitated as oxides and hydroxides with increasing pH. Oxidation-reduction potential plays an important role on the mobility of elements. Research is lacking for utilization of pH and conductance as geochemical indicator of hydrocarbon microseepage. Anomalous amounts of vanadium, chromium, nickel, cobalt, manganese mercury, copper, molybdenum, uranium, zinc, lead and zirconium are positive indicators of petroleum deposits (Duchscherer, 1984). The migrating hydrocarbons create a reducing environment in the soil and subsurface, which increases the solubility of many trace and major elements.

This paper reports the chemical alterations associated with trace metals in soils that are related to hydrocarbon microseepages above some of the major oil and gas fields of petroliferous region of Tatipaka and Pasarlapudi areas of Krishna Godavari Basin, India.

2 Geology of the study area

The Krishna-Godavari Basin is a pericratonic rift margin system with an Archean basement on the east coast of the Indian Peninsula (Kotha, 2002). It covers an area of 28,000 km^2 on land and 24,000 km^2 offshore up to 200 m deep (Gupta, 2006). The basin lies between 15°30' to 17° N latitudes and 80° to 82°30' E longitudes. Geographically, the basin lies between Kakinada in the northeast and Ongole in the southwest. Archean crystalline basement and Upper Cretaceous sedimentary outcrops demarcate the basin margin. The basin extends southeast into the deep water of the Bay of Bengal. A significant part of the onshore basinal area is covered by recent alluvium. Outcrops in the basin margin area includes the Permean Chintalapudi sandstone, Cretaceous to Jurassic Gollapalli sandstone, Raghavapuram shale and Tirupati sandstone exposed around Dwaraka Tirumala area of the West Godavari district. The Rajahmundry sandstone outcropping near the Rajahmundry and Dowleswaram areas are red, feldspathic, ferruginous and laterised sandstone of Miocene age, which is equivalent to the Ravva formation of the offshore area.

The basin is divided into the Krishna, East Godavari and West Godavari depressions separated by basement highs at Bapatla and Tanuku horsts, respectively (Rao, 1993). The East Godavari sub-basin is further divided into the Mandapeta Graben, Narsapur-Razole High and Amalapuram High. The Matsyapuri-Palakollu and Mori faults are the two major NE-SW faults. The West Godavari sub-basin is sub-divided into the Gudivada and Bantumilli graben, separated by the Kaza-Kaikaluru horst (Kumar, 1983). The tertiary alluvium sediments brought by the Godavari delta system attained

greater thickness south of the Matsyapuri-Palakollu fault as a result of continuous subsidence and growth fault related tectonics (Sanyal et al, 1998). The depositional environment varies from continental to lagoonal, littoral, infraneritic, deltaic and marine conditions. The sediments yield rich faunal assemblages like *arenaceous foraminifera* (*Ammobaculites* sp., *Ammodiscoides* sp., etc.), *Trigonia*, *Inoceramus*, *Lima*, *Pecten*, *Belemnites*, *Helicoceras*, *Cardita*, *Lamellibranchs* and *Gastropods* (Sastri et al, 1973). A geological location map of the Krishna-Godavari Basin is shown in Fig. 1.
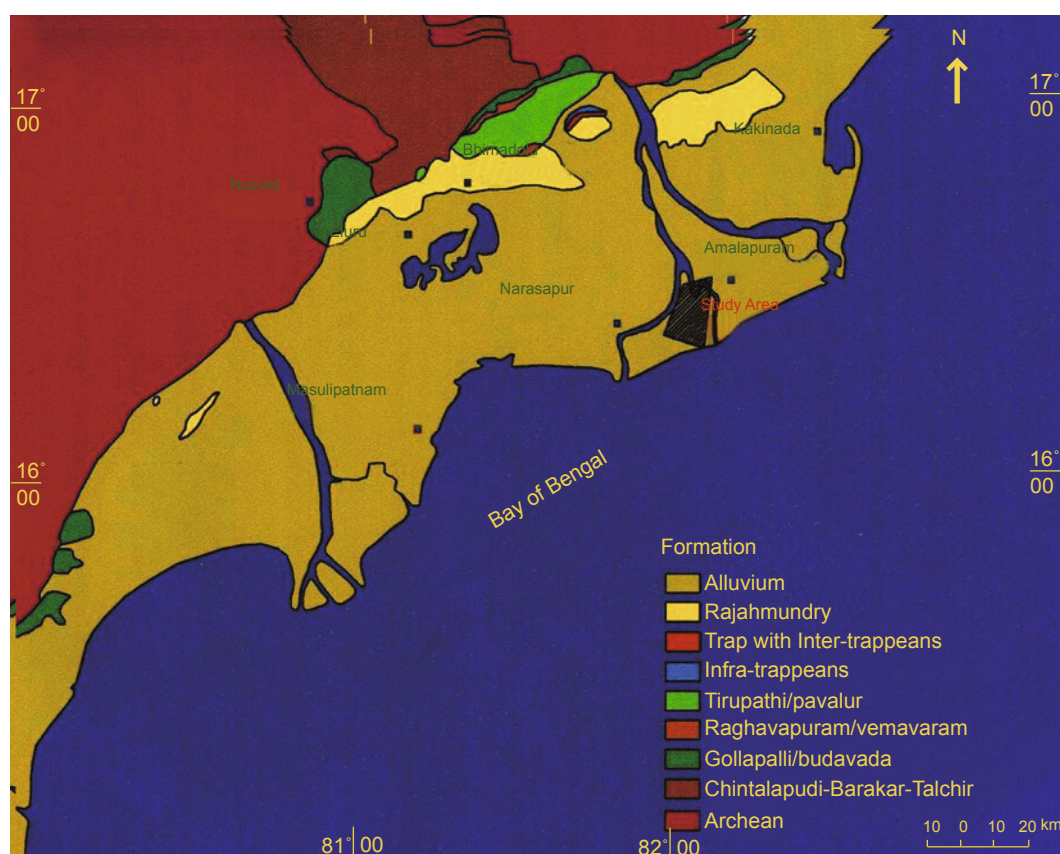

Fig. 1 Geological location map of the study area of the Krishna-Godavari Basin

2.1 Sampling

A total of 50 soil samples were collected from selected sites near oil and gas fields of the Tatipaka, Pasarlapudi areas of the Krishna Godavari Basin, Andhra Pradesh. The soil samples were collected from a depth of 2-2.5 m using a hollow metal pipe by manual hammering to the required depth. About 500 g of core soil samples collected were wrapped in aluminum foil and sealed in poly-metal packs used for analysis of adsorbed soil gases and trace element concentrations. The samples were sealed in plastic bags with their sample number and global positioning system (GPS) locations marked. Areas which were disturbed or excavated soils that were contaminated with hydrocarbons, chemicals or animal wastes and swamps & areas under water shed were avoided for sampling purposes. While collecting the samples, rocks, coarse materials, plant residues, and animal debris

have been excluded (Rasheed et al, 2008).

2.2 Light hydrocarbon analysis using gas chromatography

The light gaseous hydrocarbons were extracted from the soil samples using a gas extraction system (Horvitz, 1981). 1 gm of 63 μ particle size soil sample was used to extract light gaseous hydrocarbons after acid treatment in glass degasification apparatus and subsequent analyses on gas chromatograph (GC) and gas chromatograph–combustion–isotope ratio mass spectrometer (GC–C–IRMS) analysis for samples that had higher concentrations of hydrocarbons. During acid treatment, the dominant gas released was CO_2 which was trapped in KOH solution. The light gaseous hydrocarbons were collected by water displacement in a graduated tube fitted with rubber septa. The volume of

desorbed gas was recorded and 500 μL of desorbed gas sample was injected into gas chromatograph (Varian CP 3380) fitted with a Porapak Q column, equipped with a flame ionization detector. The gas chromatographs were calibrated using external standards with known concentrations. The quantitative estimation of light gaseous hydrocarbon constituents in each sample was made using peak area measurements and a correction for moisture content on a wet basis was also applied. The hydrocarbon concentration values of individual hydrocarbons from methane through pentane are expressed in parts per billion (μg/kg).

2.3 Physico-chemical analysis

The physico-chemical properties, such as pH, electrical conductivity (EC), total dissolved solids (TDS) and oxidation-reduction potential (Eh) in soil samples, were analyzed on pH/EC/Eh Bench top meter (Thermo Electron Corp. Orion 5 star).

2.4 Analysis of trace elements using inductively coupled plasma mass spectrometer (ICP-MS)

The analysis of trace elements in the soil samples was carried out using ICP–MS (Perkin-Elmer Sciex Elan DRC II) at the National Geophysical Research Institute, Hyderabad. The chemical analyses of the soil samples were carried out using ICP–MS (Perkin-Elmer Sciex Elan DRC II). An ICP–MS is a type of mass spectrometry based on coupling together an inductively coupled plasma as a method of producing ions (ionizations) with a mass spectrometer as a method of separating and detecting ions. ICP–MS is highly sensitive and capable of determination of a range of metals and several non-metals at concentrations that are less than one part in 10^{12}.

The soil samples were dried and sieved using a 63 μ sieve. A sample size of 0.05 g was taken in a Teflon beaker. 10 mL of acid mixture (7:3:1) comprising of HF, HNO_3 and $HClO_4$ was added to the sample and left over night at room temperature, and the samples were dried at 200 ℃ on a hotplate. 20 mL of 1:1 HNO_3 and double distilled water were added and heated for 10 minutes on a hotplate at 200 ℃. Finally, rhodium as internal standard was added and made up to 200 mL with distilled water. A precision of 1%-7% relative to standard deviation and comparable accuracy was obtained by this method. The analytical procedures and protocols followed at NGRI are the same as given in Naqvi et al (2006), Balaram et al (1996) and Rao and Govil (1995).

3 Results and discussion

3.1 pH

The pH of the soils in the study area ranges from 5.72 to 8.88 with an arithmetic mean of 8.18 and standard deviation of 0.58 (Table 1). Most of the soil samples are alkaline in nature. It is observed that, the pH ranges between 6 and 8 above the oil and gas accumulations, and it is below 6 around the periphery. The acidic nature of soils prevents adsorption of hydrocarbons (Horvitz, 1985). The highly acidic nature

around the periphery involves in the mobilization of certain major and minor and trace elements to the oxidizing areas and precipitation as oxides and hydroxides. The concentration of certain elements above oil field areas and depletion in the soils around the periphery is the main cause of acidity of the soils. The soil acidity is the main reason for creating an oxidation-reduction potential or front. The pH is observed to be 7 to 8 around gas producing areas; around oil producers, it is 8 to 8.5 and around non-producing oil/gas areas, it is greater than 8.5 (Rao, 2006). The pHs in the present study area are observed to be in the range for a gas producing area, which is true as the majority of the wells in the KG Basin are gas producers.

Table 1 The statistical analysis of physico-chemical properties in soil samples of the study area

	pH	EC μS/cm	Eh (ORP) mV
Minimum	5.72	76	142
Maximum	8.88	3490	295
Arithmetic mean	8.18	888	205
Median	8.38	417.5	207
Standard deviation	0.58	963	32

3.2 EC

The amount of soluble salts in soils is generally estimated by measuring the electrical conductivity of aqueous soil extracts. The EC of the near surface soils in the study area ranges from 76 to 3,490 μS/cm at 25 ℃, with an arithmetic mean of 888 μS/cm and a standard deviation of 963 μS/cm (Table 1). The specific conductance is a supporting parameter for recognizing gas microseepages. Electrical conductivity is found to be a useful parameter to locate gas and gas associated oil reservoirs. There is a recognizable increase in electrical conductivity above areas of gas accumulation (Heemstra et al, 1979), but little effect over oil and barren areas.

3.3 Eh

Petroleum microseepage causes several chemical reactions in the near surface soil. The near surface change of Eh and pH conditions in response to light hydrocarbon microseepage and its microbial oxidation drives a localized redistribution of trace elements. Oxidation-reduction potential plays an important role in the mobility of elements. The Eh in the near surface soils of the study area ranged from 142 to 295 mV with arithmetic mean of 205 mV and standard deviation of 32 mV (Table 1). Eh measures the tendency of an environment to oxidize or reduce substrates and the values decrease from oxidizing to reducing environments (Maier et al, 2009).

Petroleum seeps into the soil substrate altering the existing Eh and pH conditions, and causes a reducing environment, where solubility increases and transportation of trace metals occurs. The reducing environment created

by the petroleum oxidizes and mobilizes various elements. The most common carbonates cannot be oxidized, $CaCO_3$ and $MgCaCO_3$ for instance are fully oxidized and cannot be further oxidized. However, siderite can be oxidized. Chemical solubility is a prerequisite for physical mobility of the trace metals (Robinson et al, 2005). Trace metals become more mobile under acidic and reducing conditions as long as slightly soluble metal sulfides do not precipitate (Rennert and Mansfeldt, 2006). Given a steady pH, even the reducing conditions enhance the mobility of trace metals (Charlatchka and Cambier, 2000; Chuan et al, 1996). The hydrodynamic flow in the soil moves the solubilized elements around although they do not necessarily leave the reduced area. Concentration into the "anomalous" areas occurs where reducing groundwater is low in pH and Eh and intermingles with groundwater of higher Eh and pH. This results in deposition of carbonate, sulfide, phosphate, and oxide minerals, which incorporate the trace elements into the compounds (Tedesco, 1995).

3.4 Trace element geochemistry

Trace element concentrations (mg/kg) in soil samples of the study area are given in Table 2. The trace metals scandium (Sc), vanadium (V), chromium (Cr), cobalt (Co), nickel (Ni), copper (Cu), zinc (Zn), barium (Ba) and strontium (Sr) are considered for the study and the concentrations of each of the trace elements varied in the following manner. The concentrations of each of the trace elements Sc (8 to 40 mg/kg), V (197 to 489 mg/kg), Cr (106 to 287 mg/kg), Co (31 to 52 mg/kg), Ni (65 to 110 mg/kg), Cu (88 to 131 mg/kg), Zn (88 to 471 mg/kg), Ba (263 to 3,091 mg/kg) and Sr (119 to 218 mg/kg). The concentration of trace metals along with normal values in soil samples of the study area is given in Table 3. The concentrations of trace elements in normal soils are as follows for Sc (7 mg/kg), V (90 mg/kg), Cr (70 mg/kg), Co (8 mg/kg), Ni (50 mg/kg), Cu (30 mg/kg), Zn (90 mg/kg), Ba (500 mg/kg) and Sr (250 mg/kg) (Bowen, 1979). It was observed that the concentrations of trace elements were tremendously increased when compared with normal concentrations in soils.

The concentration distribution maps of these trace elements in the study area can be seen with their respective composite anomaly maps with ΣC_{2+} (Figs. 2, 3, 4, 5, 6, 7, 8 and 9). The trace elements anomalies are located in the Tatipaka and Pasarlapudi areas and west of Pasarlapudi lanka. Trace elements concentration distribution showed good correlation with the existing oil and gas wells. Increased amounts of soluble Ni, V, Cu, Cr, Zn and Co have been observed in the reducing environment caused by the seepage of hydrocarbons. In the reducing environment, solubility of metals increases and transport occurs. Due to hydrodynamic flow in the soil, the solubilized elements move around, although they do not leave the reducing area (Tedesco, 1995). When oil migrated, the trace-elements within the oil also migrate, and these can get added to trace metals of soil and they may experience the same chemical changes as those of soil trace metals in the hydrocarbon seepage areas. Vanadium and nickel are the only heavy metal elements that occur

in crude oils in appreciable quantities as soluble organic complexes (Al-Shahristani and Al-Atyia, 1972). Other trace elements such as Mn, Cr, Cu, Zn, Sb, Se, Br, I anomalies related to hydrocarbon reservoirs have been reported by Clark et al (2003). Because a boundary is formed between the reducing and oxidizing zone by the deposition of carbonate, oxide and sulphide minerals, several metals including Ni, V, Cu, Cr, Zn and Co are mobilized in soils and accumulated around the hydrocarbon anomaly when reducing conditions are created due to hydrocarbon accumulation (Nissenbaum and Swaine, 1976). Recent geochemical studies carried in Neuquen Basin of Argentina have revealed that the trace metal vanadium is transported to the surface by the uplifting of microbubbles of hydrocarbon gases. This is known from the fact that vanadium content is closely related to bacterial anomalies that feed on the seeping hydrocarbons. It is reported that the concentrations of V, Zn and Ba are seen concentrated over the faults and lineaments, which might have provided the migrational pathway for their transportation (Larriestra et al, 2010). Hence, the concentrations of trace elements in the present study area, were found to be extremely high, which is evidence of hydrocarbon microseepage in the Krishna Godavari Basin.

3.5 Integration with adsorbed soil gases

Trace elements concentrations have been used as indirect indicator for hydrocarbon microseepage (Madhavi et al, 2011). This study compares the trace elements with direct detection of adsorbed light hydrocarbon soil gas concentrations and explains how the trace elements concentrations vary near hydrocarbon anomalies. The compositional correlation is fundamental to understanding the link between trace element concentrations and hydrocarbon anomalies. The adsorbed soil gas analysis showed the presence of high concentrations of ΣC_{2+} (C_2H_6, C_3H_8 and n-C_4H_{10}) ranged between 7 to 222 µg/kg. Integrated maps of trace elements over adsorbed light gaseous hydrocarbons (ΣC_{2+}) anomalies showed good correlation with the existing oil and gas wells (Figs. 2, 3, 4, 5, 6, 7, 8 and 9). The adsorbed soil gas anomalies were seen in the centre and also adjacent to the trace metal anomalies. Most of the trace metal anomalies (Cr, V, Co and Zn) were observed in north-east and south-west parts of the study area, and scanty in the middle part of the study area except for Cu and Ni. By observing the composite maps of hydrocarbons and trace elements, clearly indicate that trace elements are accumulated around the hydrocarbon anomalies, creating a halo pattern. The study shows that most of the trace elements follow a halo pattern encircling an apical pattern of hydrocarbon seepage indicating the presence of a reduced body in the subsurface. The composite maps of hydrocarbons and trace elements, clearly indicate that trace elements are accumulated around the hydrocarbon anomalies, creating a halo pattern. Trace elements occur as a near surface 'halo' which will be pathfinders or otherwise important in the search for hydrocarbon anomalies (Siegel, 1974). Duchscherer (1983) reported that trace element associations form organometallic compounds, are found "haloed" or concentrated over or around underlying hydrocarbon reservoirs. In the present

Table 2 Trace element concentrations (mg/kg) in soil samples of the study area

Sample ID	Sc	V	Cr	Co	Ni	Cu	Zn	Ba	Sr
KGW-01	26	250	140	40	91	118	114	340	151
KGW-02	25	256	164	40	93	120	146	391	145
KGW-03	19	197	107	31	65	88	89	264	133
KGW-04	25	380	216	50	89	118	164	3092	173
KGW-05	19	265	151	40	73	96	117	449	218
KGW-06	23	278	175	45	83	104	251	750	193
KGW-07	20	272	129	45	79	106	118	481	187
KGW-08	27	280	173	47	99	130	129	416	137
KGW-09	40	490	288	52	110	102	471	378	131
KGW-10	29	323	198	41	90	131	262	1058	119
KGW-11	25	273	168	42	93	121	132	707	166
KGW-12	31	270	193	38	106	127	130	557	126
KGW-13	8	280	154	40	104	131	127	372	129
KGW-14	28	268	185	41	102	130	142	543	123
KGW-15	26	250	140	40	91	118	114	340	151
KGW-16	29	279	129	40	99	130	142	378	193
KGW-17	23	271	198	45	93	131	114	340	123
KGW-18	26	273	193	38	90	127	130	481	166
KGW-19	20	489	288	42	93	118	117	557	119
KGW-20	19	269	185	31	104	131	262	1058	145
KGW-21	31	197	164	41	79	121	89	543	131
KGW-22	28	268	164	40	93	120	146	372	151
KGW-23	25	265	151	40	91	130	471	391	187
KGW-24	25	255	107	41	89	104	251	750	218
KGW-25	8	380	140	40	65	106	127	391	173
KGW-26	19	277	216	50	102	118	164	449	137
KGW-27	25	255	168	40	83	88	132	416	129
KGW-28	27	279	175	45	106	120	146	3092	133
KGW-29	25	249	154	47	110	96	118	264	145
KGW-30	39	323	173	52	73	102	129	707	126
KGW-31	26	279	216	41	79	118	127	481	119
KGW-32	24	269	151	40	99	96	118	3092	145
KGW-33	22	249	154	31	65	131	129	1058	133
KGW-34	25	279	164	40	93	102	471	264	187
KGW-35	39	277	173	45	65	88	146	750	166
KGW-36	18	197	151	50	89	130	129	449	218
KGW-37	18	380	107	47	99	104	251	3092	173
KGW-38	26	265	129	50	102	130	114	264	193
KGW-39	19	271	129	45	83	127	118	543	218
KGW-40	28	380	175	40	91	106	89	750	131
KGW-41	18	255	107	40	106	88	164	391	133
KGW-42	19	273	185	40	90	118	89	481	137
KGW-43	18	265	175	45	73	96	117	416	137
KGW-44	26	279	193	41	83	106	251	416	173
KGW-45	22	277	140	31	73	121	130	372	193
KGW-46	24	489	168	45	79	118	262	378	126
KGW-47	2	323	288	38	93	104	142	340	123
KGW-48	8	197	173	42	104	130	132	707	151
KGW-49	31	268	198	47	89	131	117	449	187
KGW-50	25	271	216	52	110	120	164	557	129

Table 3 Descriptive statistics of trace elements (mg/kg) in soil samples of the study area

Elements	Minimum	Maximum	Average	Standard deviation	Median	Normal soil values*
Sc	8	40	24	7	25	7
V	197	490	288	65	272	90
Cr	107	288	171	41	168	70
Co	31	52	42	5	41	8
Ni	65	110	90	12	91	50
Cu	88	131	115	14	118	30
Zn	89	471	166	91	130	90
Ba	264	3092	712	735	449	500
Sr	119	218	155	30	145	250

Notes: * Values from Bowen (1979)
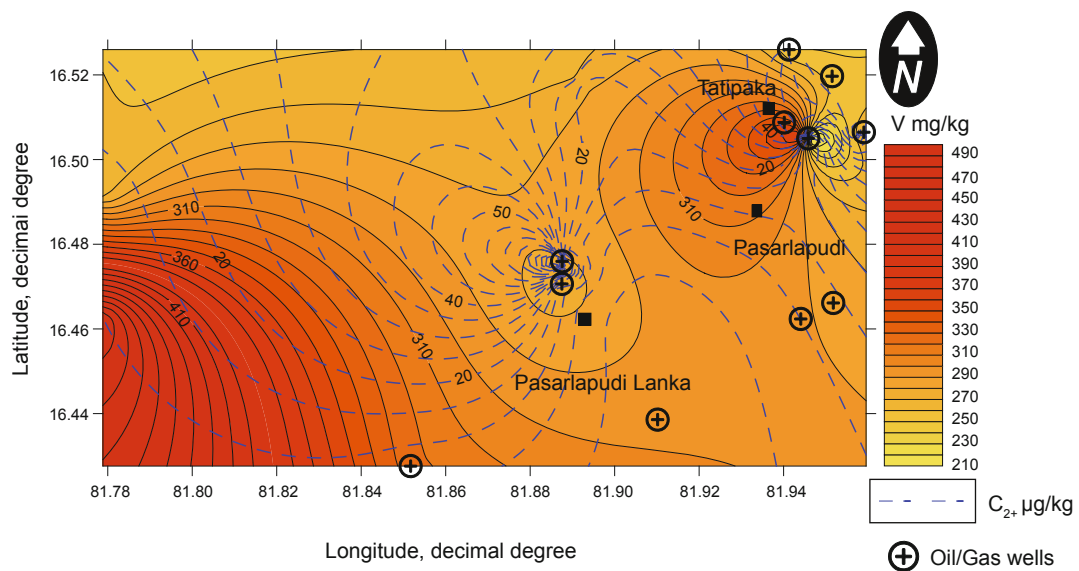
Fig. 2 Composite map of vanadium and adsorbed soil gas concentration in oil and gas fields of the study area
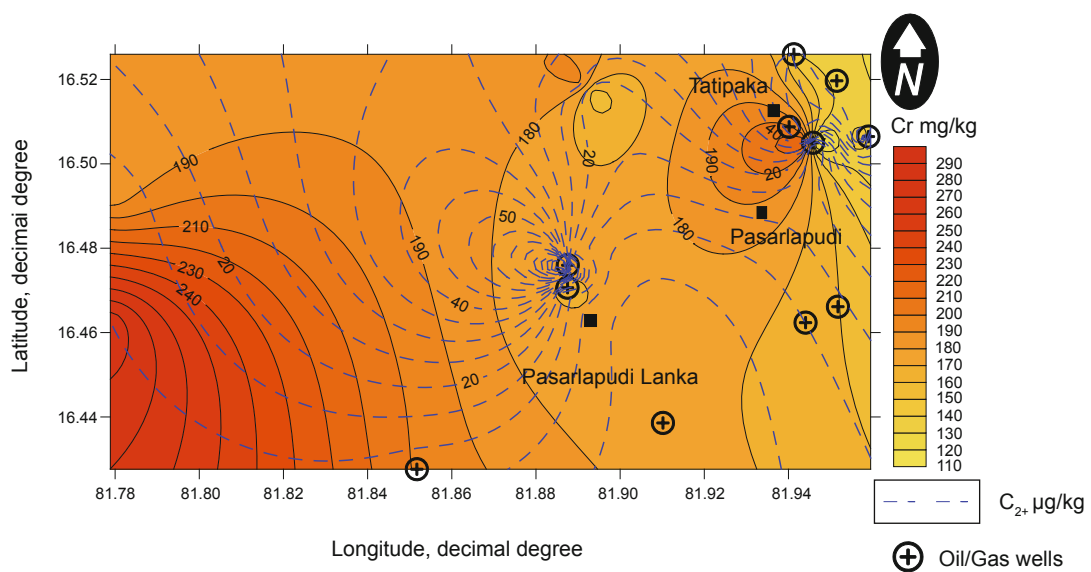
Fig. 3 Composite map of chromium and adsorbed soil gas concentration in oil and gas fields of the study area
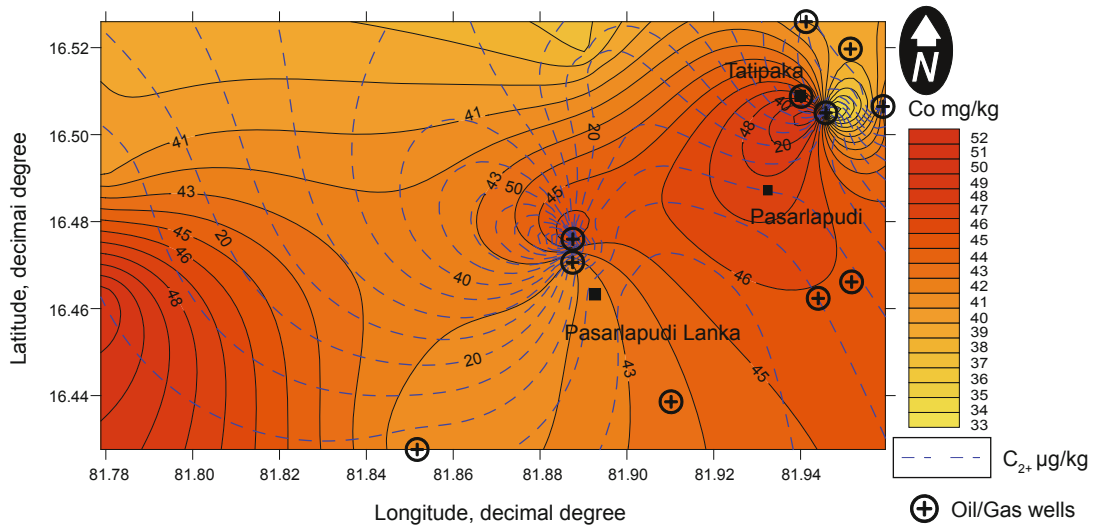
Fig. 4 Composite map of cobalt and adsorbed soil gas concentration in oil and gas fields of the study area
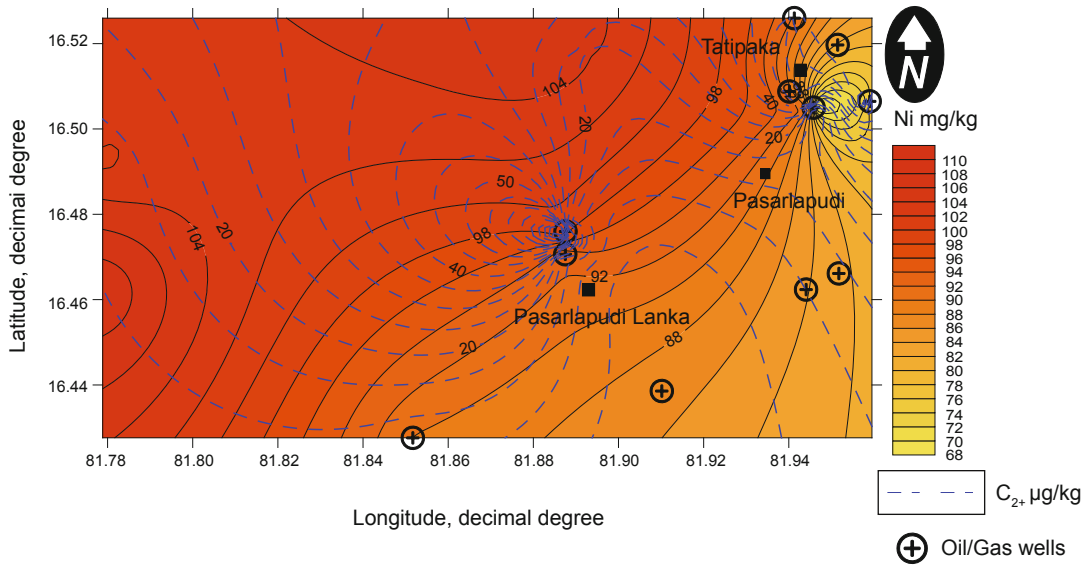
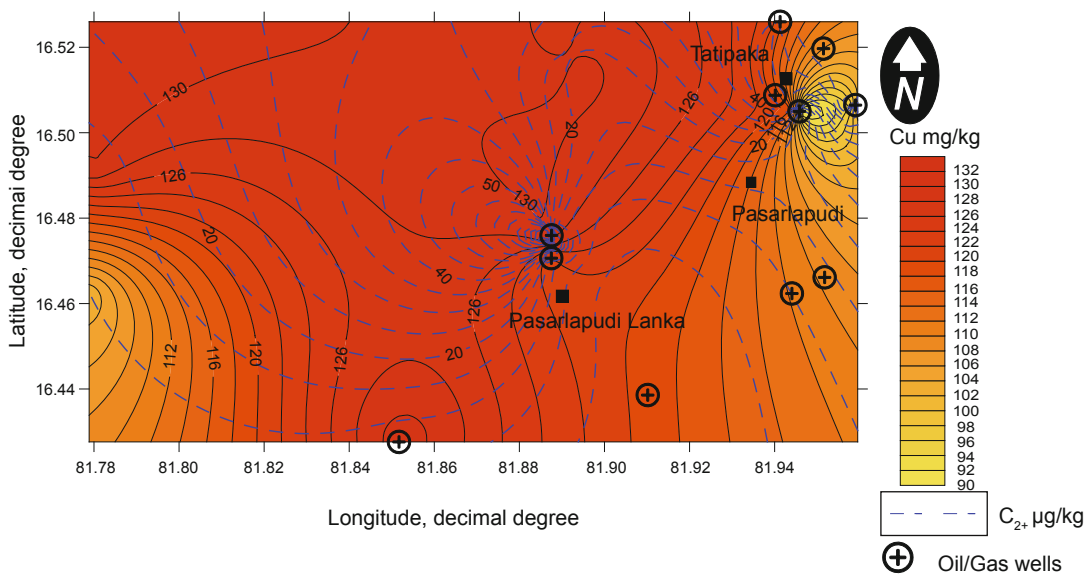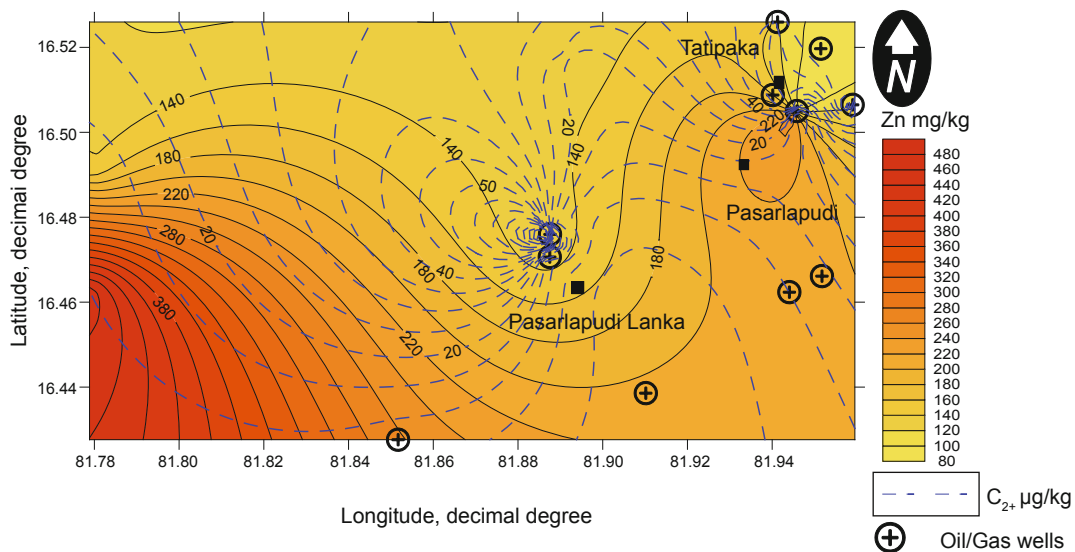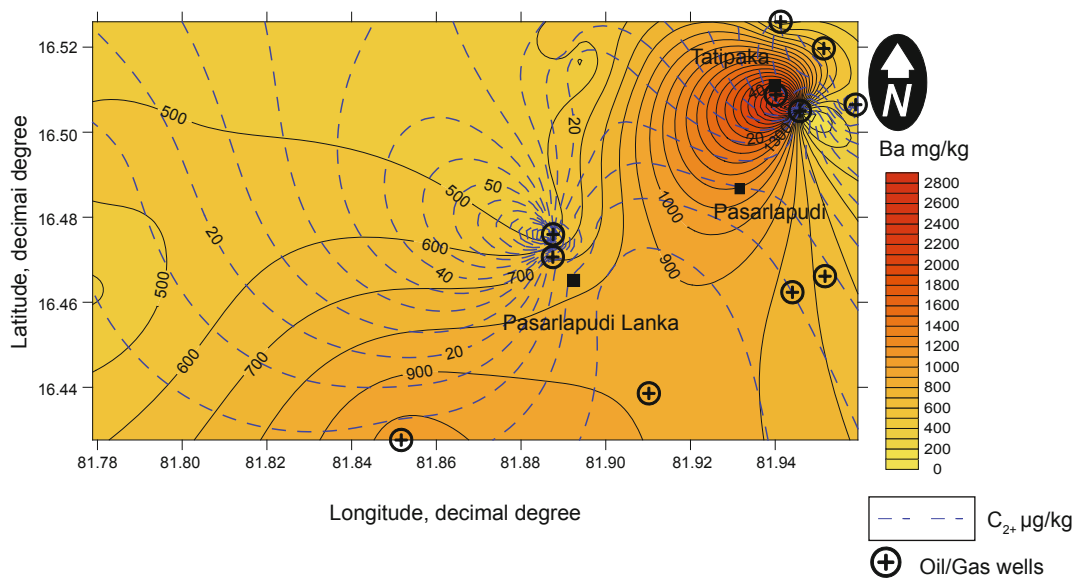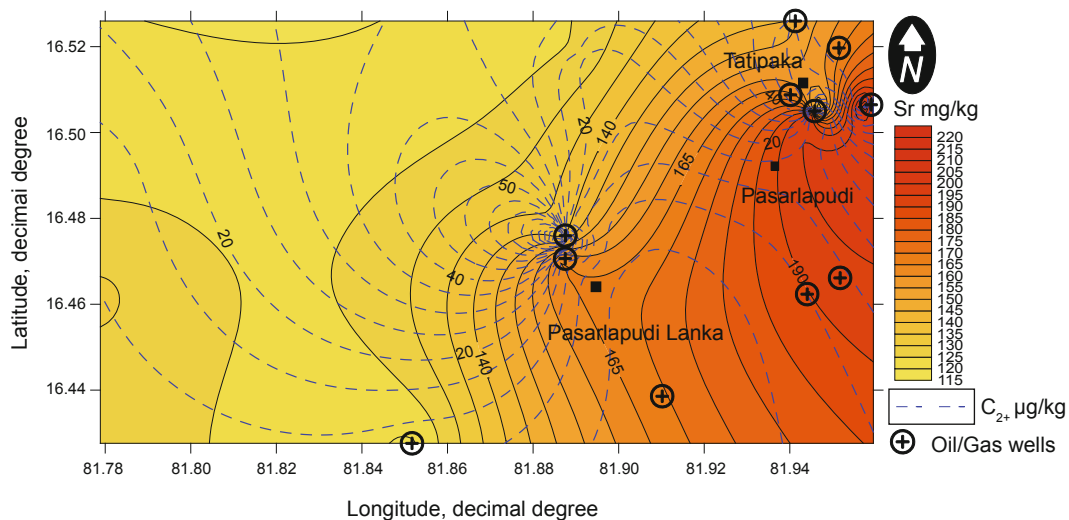**Fig. 5** Composite map of nickel and adsorbed soil gas concentration in oil and gas fields of the study area

Fig. 6 Composite map of copper and adsorbed soil gas concentration in oil and gas fields of the study area

Fig. 7 Composite map of zinc and adsorbed soil gas concentration in oil and gas fields of the study area

Fig. 8 Composite map of barium and adsorbed soil gas concentration in oil and gas fields of the study area

Fig. 9 Composite map of strontium and adsorbed soil gas concentration in oil and gas fields of the study area

study area of the Krishna Godavari Basin, the concentration distribution maps of vanadium, nickel, chromium, iron, cobalt, copper and manganese showed inverse correlation whereas those of strontium and barium showed positive correlation with the migrating hydrocarbons.

The carbon isotopic composition of $\delta^{13}C_1$ of the samples ranges between $-36.6‰$ to $-22.7‰$ (PeeDee Belemnite) values indicate thermogenic origin, which presents convincing evidence that the adsorbed soil gases collected from these sediments are of catagenetic origin. Enrichment of trace elements seen around the hydrocarbon anomaly helps verify the correlation between microseepage and trace elements concentrations. That enrichment of trace elements in one or more prospective zones around hydrocarbon anomalies can suggest that trace elements may be a pathfinder or otherwise an important search tool for hydrocarbon anomalies.

4 Conclusions

The concentrations of trace elements showed excellent correlation with existing oil and gas wells in the study area. The anomalies of trace elements are found to be in good agreement with geological information available for the area and correlated with the major oil and gas fields of the Krishna Godavari Basin. The composite anomaly maps of ΣC_{2+} and trace elements correlate well with the underlying hydrocarbon reservoirs in the study area. The increase in the concentrations of trace metals near oil/gas producing areas, suggests a soil chemical change to a reducing environment, presumably due to the influence of hydrocarbon microseepage, which could be applied with other geo-scientific data to identify areas of future hydrocarbon exploration in the frontier areas. The trace element studies in correlation with adsorbed soil gas studies confirmed the seepage of lighter hydrocarbon accumulations from the subsurface and help in finding hydrocarbon microseepage in petroliferous regions. The experimental research work has the potential to reduce risk in petroleum exploration. This is of high significance, especially now when globally petroleum reserves replacement is becoming a challenge.

Acknowledgements

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(Edited by Hao Jie)